

# An Analysis on the Retrieval of Raindrop Size Distribution by Dual-Frequency Precipitation Radar

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## Abstract

The retrieval of raindrop size distribution has a great significance for the study on the microphysical process of cloud and precipitation, as well as the formation and development of precipitation systems. In this paper, given raindrop size distributions, the T-matrix method is used to simulate the radar detection parameters for solid, mixed and liquid particles, with which the DPR algorithm is used to retrieve the raindrop size distribution to analyze the effect and the sensitivity of the algorithm for different phases. The results show that, for snow particles, when the normalized scale parameter  $N_w$  is fixed, the retrieved median diameter  $D_0$  is less than the truth value, and the retrieval error increases with the increase of  $D_0$ ; When  $D_0$  is fixed, the retrieved  $N_w$  is larger than the true value and they have a notable linear relationship. With the increase of the  $N_w$ , the retrieval error decreases continuously while the retrieved value changes little. For the rain particles, the retrieval of  $D_0$  is very good, and the retrieval of  $N_w$  is slightly worse. For the mixed particles, the retrieve  $D_0$  is less than the true value, and the retrieve  $N_w$  is larger than the true value. The retrieval accuracy increases with the decrease of the mixing ratio.

**Keywords:** Drop size distribution; Dual-frequency radar; Retrieval

**Abbreviations:** DSD: Drop Size Distribution; PR: Precipitation Radar; GPM: Global Precipitation Measurement; DPR: Dual Frequency Precipitation Radar

## Introduction

In the research field of radar meteorology, it is important to obtain the rain Drop Size Distribution (DSD) information. The retrieval of DSD is of great significance for the study on the microphysical process of cloud and precipitation, and on the formation and development of precipitation systems. It can be also used to analyze the characteristics of precipitation distribution and to improve the accuracy of radar precipitation estimation. The space borne radar has a satellite platform of a high orbit and can provide global precipitation observation. The PR (Precipitation Radar) carried by the TRMM (Tropical Rainfall Measuring Mission) is the first spaceborne precipitation radar. The Global Precipitation Measurement (GPM) is the successor to TRMM, and the Dual Frequency Precipitation Radar (DPR) onboard GPM uses dual frequency technique to detect rainfall.

Ulbrich [1] found that the three parameters of the Gamma distribution (normalized scale factor  $N_w$ , shape parameter  $\mu$ , and median drop diameter  $D_0$ ) were not independent of each other and had high Pearson correlation coefficients among them. Zhang et al. [2] calculated the other two parameters of Gamma distribution by the relationship between the shape parameters  $\mu$  and  $\Lambda$ . Seto et al. [3] developed an algorithm to retrieve DSD based on spaceborne dual-frequency radar, namely the DFR (Dual-Frequency Reflectivity Ratio) method. Chandrasekar et al. [4] showed that DFR in rain area usually encountered the "double

value” problem. In the current DPR algorithm, the relationship between precipitation rate  $R$  and  $D_m$  is adjusted by one or several adjustment factors to eliminate the “double value” problem when the precipitation rate is small. Liao et al. [5] evaluated the advantages of dual-frequency algorithm over single-frequency algorithm in current DPR algorithms. In this paper, the T-matrix method is used to simulate the radar detection parameters of solid, mixed and liquid particles when raindrop size distribution is given, and then the DPR algorithm is used to retrieve the DSD to analyze the effect of the algorithm for different phases.

## Basic Theory

### Particle model

In the study of this paper, rain particles are pure liquid particles, mixed particles are uniformly mixed particles of ice, air and water, and solid particles are uniformly mixed particles of ice and air. The particle density is calculated as in Bohren et al. [6]:

$$\rho_b = P_w \rho_w + P_i \rho_i + P_a \rho_a \quad (1)$$

Where,  $P_w$ ,  $P_i$  and  $P_a$  respectively represent the percentage of water, ice and air in the whole particle  $\rho_w$ ,  $\rho_i$  and  $\rho_a$  respectively represent the densities of water, ice and air, which are fixed as  $1\text{g}/\text{cm}^3$ ,  $0.92\text{g}/\text{cm}^3$  and  $0\text{g}/\text{cm}^3$ . In this paper, the shape of precipitation particles is oblate ellipsoid, and the relationship between the axial ratio of the oblate ellipsoid and the equivalent diameter is modeled by Brandes et al. [7]. The specific calculation formula is:

$$r = 0.9951 + 0.02510D - 0.3644D^2 + 0.005030D^3 - 0.0002492D^4 \quad (2)$$

where  $r$  is the axial ratio of the oblate ellipsoid, and  $D$  is the equivalent diameter of the particle. In this paper, it is assumed that: 1) The dual-frequency (Ku and Ka band) radar observes stratiform precipitation; 2) The temperature of the liquid layer is  $10^\circ\text{C}$ ; 3) The solid particle density is  $0.2\text{g}/\text{cm}^3$ ; 4) The tilt angles of the particles  $\theta$  satisfy the Gaussian distribution with a mean value of  $0$  and a variance of  $6^\circ$ ; 5) The rotation angles of the particles  $\Phi$  satisfies satisfy an average distribution of  $0^\circ\sim 180^\circ$ ; 6) The particle diameter is  $0\sim 8\text{mm}$ .

### DSD model

Ulbrich proposed the Gamma distribution to describe the characteristics of DSD [1]. The Gamma distribution includes the concentration parameter  $N_0$  ( $\text{m}^{-3}\text{mm}^{-1}$ ), the shape factor  $\mu$  (dimensionless) and the scale parameter  $\Lambda$  ( $\text{mm}^{-1}$ ):

$$N(D) = N_0 D^\mu \exp(-\Lambda D) \quad (3)$$

where  $D$  is the equivalent volume diameter (mm). In the Gamma distribution, since the unit of  $N_0$  is not consistent with the DSD concentration  $N(D)$  and has no clear physical meaning, it is not conducive to the retrieval of concentration parameters. Willis proposed a “normalized” Gamma distribution as shown in (4) [8], which introduces the volume median diameter  $D_0$  (mm) and normalized scale factor  $N_w$  ( $\text{m}^{-3}\text{mm}^{-1}$ ):

$$N(D) = N_w f(\mu) \left(\frac{D}{D_0}\right)^\mu \exp\left[-(3.67+\mu)\frac{D}{D_0}\right] \quad (4)$$

$$N_w = \frac{N_0}{f(\mu)} D_0^\mu \quad (5)$$

$$f(\mu) = \frac{6}{3.67^4} \frac{(3.67+\mu)^{\mu+4}}{\Gamma(\mu+4)} \quad (6)$$

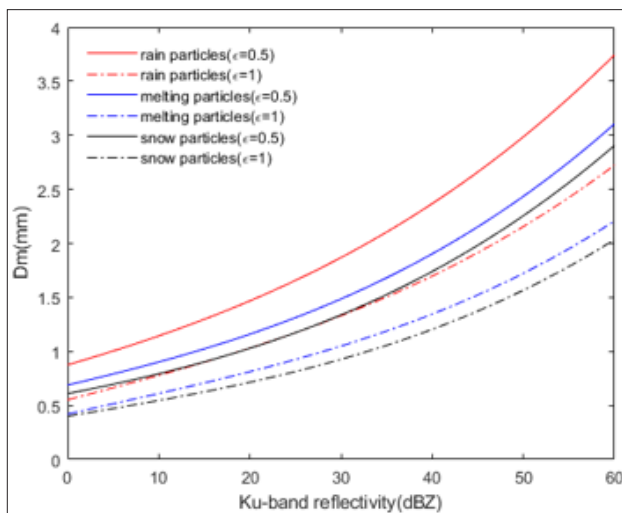
In this paper, according to the DPR algorithm,  $\mu$  is set to  $0$  in the ice and melting ice regions and set to  $3$  in the rain region.

### The retrieval algorithm of DSD

The current DPR algorithm adds  $R$ - $D_m$  relations to constrain the retrieval of DSD. If the Gamma distribution has a fixed shape factor  $\mu$  or a mass-weighted diameter  $D_m$  function, it will have only one parameter  $D_m$ .

$$R = \varepsilon^r p D_m^q \quad (7)$$

When the DSD satisfies the Gamma distribution, and  $\mu \geq 2$ , there is a relationship between  $D_m$  and  $D_0$ :  $(3.67+\mu)D_m = (4+\mu)D_0$ . According to the analysis of observations by Kozu et al. [9] the values of  $r, p, q$  in the current algorithm are  $0.401$ ,  $6.131$  and  $4.649$  in stratiform precipitation, and  $1.370$ ,  $5.420$  and  $4.258$  in convective precipitation, respectively [9].  $\varepsilon$  ranges from  $0.2$  to  $5$ , and its optimal value is determined by the minimization of the difference between the simulated and the measured radar reflectivity, and between the calculated and the estimated PIA (the integral attenuation caused by precipitation particles from the surface to the radar).  $Z_{Ku}$  and  $D_m$  are related in a one-to-one pattern when  $\varepsilon$  is given so that  $D_m$  can be uniquely derived from  $Z_{Ku}$ , and then  $N_w$  can be obtained from  $D_m$  and  $Z_{Ku}$ . It can be seen from (Figure 1) that for the same  $Z_{Ku}$ ,  $D_m$  (for rain particles)  $> D_m$  (for melting particles)  $> D_m$  (for snow particles), and  $D_m$  decreases as  $\varepsilon$  increases.

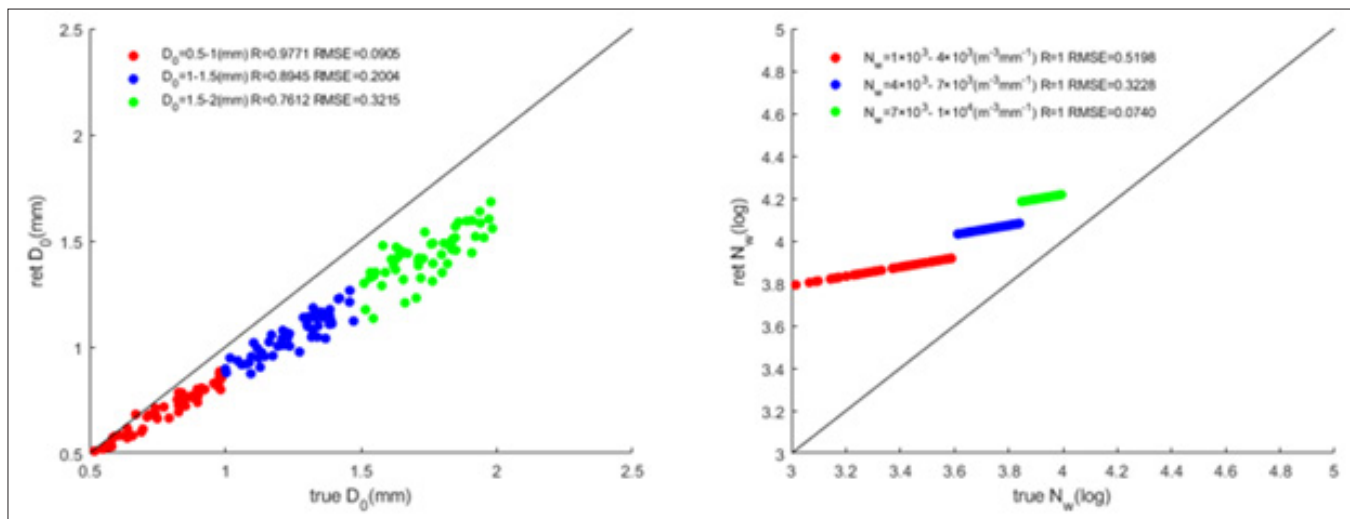


**Figure 1:**  $D_m - Z_{Ku}$  relations for  $\epsilon$  value of 0.5 (solid lines) and 1 (dashed lines) for rain particles (red), melting particles (blue), snow particles (black) in the case of stratiform precipitation.

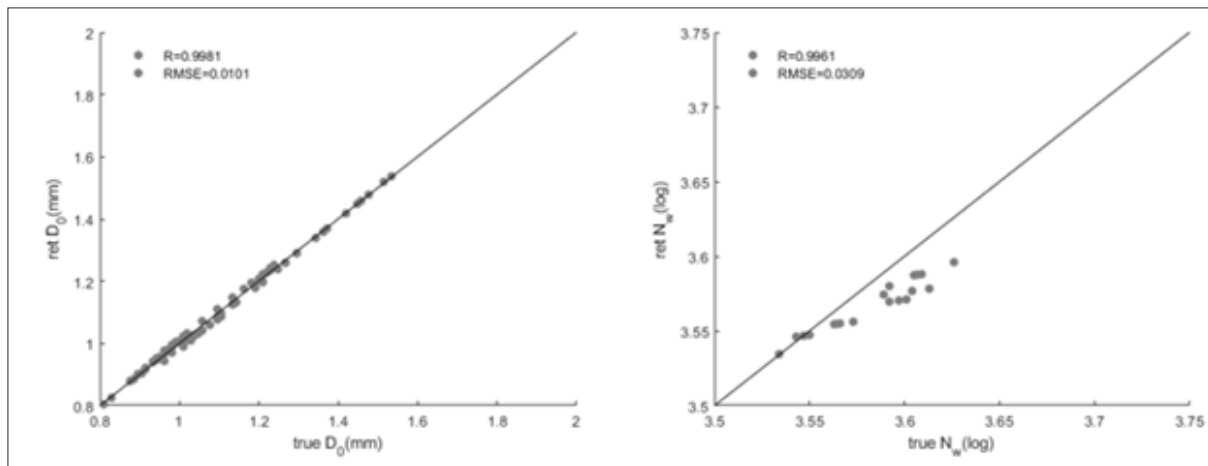
**The Performance of the DPR Retrieval Algorithm**

Because the values of  $N_w$  usually have a very large range, they are shown by the logarithmic scale in this study. For the snow particles, As shown in Figure 2, when  $N_w$  is fixed and the range of  $D_0$  is 0.5mm~1.0mm, 1.0mm~1.5mm and 1.5mm~2.0mm, the retrieved  $D_0$  is less than the truth value and the retrieval error increases with the increase of  $D_0$ ; When  $D_0$  is fixed and the range of  $N_w$  is  $1 \times 10^3 m^{-3} mm^{-1} \sim 4 \times 10^3 m^{-3} mm^{-1}$ ,  $4 \times 10^3 m^{-3} mm^{-1} \sim 7 \times 10^3 m^{-3} mm^{-1}$  and  $7 \times 10^3 m^{-3} mm^{-1} \sim 1 \times 10^4 m^{-3} mm^{-1}$ , the retrieved  $N_w$  is larger than the true value and they have a notable linear relationship. With

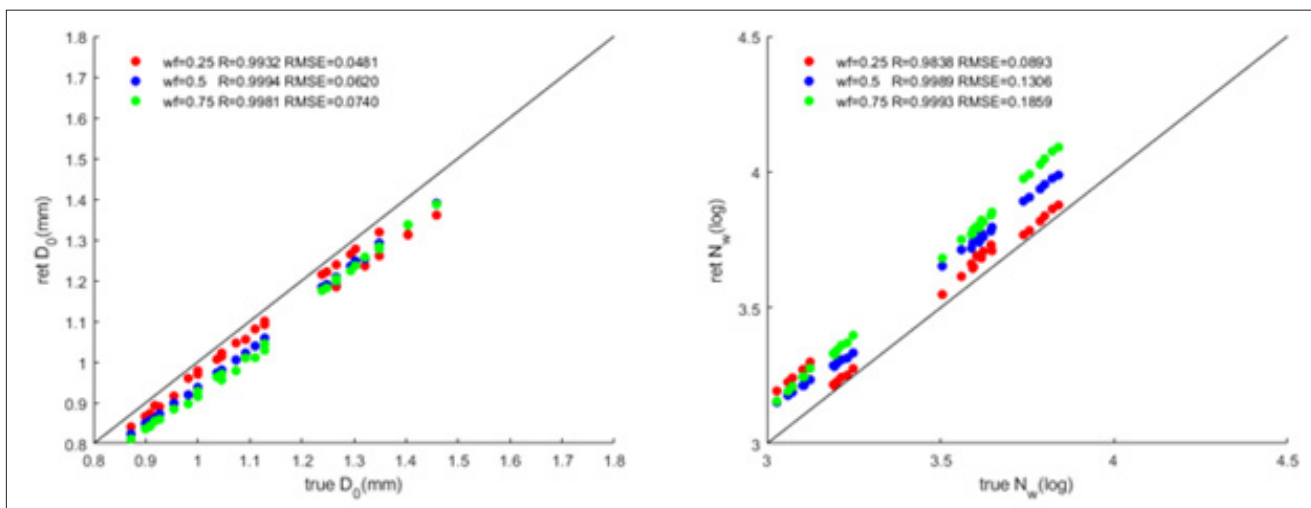
the increase of  $N_w$  the retrieval error decreases while the retrieved value changes little. For rain particles, the DSD parameters in DPR product are taken as the true value. It can be seen from (Figure 3) that the retrieval of  $D_0$  is very good and the retrieval of  $N_w$  is slightly worse. For the mixed particles, the DSD parameters in DPR product are taken as the true value, and the mixing ratio is set to 0.25, 0.5 and 0.75, respectively. As shown in Figure 4, the retrieved  $D_0$  is less than the true value, and the retrieved  $N_w$  is larger than the true value. The retrieval accuracy increases with the decrease of the mixing ratio.



**Figure 2:** The true values of  $D_0$  (left) and  $N_w$  (right and in log scale) and retrieved values by the DPR algorithm for the snow particles.



**Figure 3:** The true values of  $D_0$  (left) and  $N_w$  (right and in log scale) and the retrieved values by the DPR algorithm for the rain particles.



**Figure 4:** The true values of  $D_0$  (left) and  $N_w$  (right and in log scale) and the retrieved values by the DPR algorithm for the melting particles with the water fraction of 0.25 (red), 0.5 (blue), 0.75 (green).

### Summary and conclusion

In this paper, the T-matrix method is used to simulate the radar detection parameters of solid, mixed and liquid particles given the DSD, and then the DPR algorithm is used to retrieve the DSD based on the radar detection parameters. The retrieval results are analyzed by comparison with the true values and the conclusions are as follows:

a. For the snow particles, when  $N_w$  is fixed, the retrieved  $D_0$  is less than the truth value, and the retrieval error increases with the increase of  $D_0$ ; When  $D_0$  is fixed, the retrieved  $N_w$  is larger than the true value, and the retrieved value and the true value have a notable linear trend. With the increase of the  $N_w$  true value range, the retrieval error decreases while the retrieved value changes little.

b. For the rain particles, the retrieval of  $D_0$  is very good, and the retrieval of the  $N_w$  retrieve is slightly worse.

c. For the mixed particles, the retrieved  $D_0$  is less than the true value, and the retrieve  $N_w$  is larger than the true value. The retrieval accuracy increases with the decrease of the mixing ratio.

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